

**The contribution of the  
fine sediment fraction  
to the Fluffy Layer  
Suspended Matter  
(FLSM)\***

OCEANOLOGIA, 44 (4), 2002.  
pp. 513–527.

© 2002, by Institute of  
Oceanology PAS.

**KEYWORDS**

Lead 210  
Polonium 210  
Isotopic disequilibria  
Pomeranian Bay  
Baltic Sea

JANUSZ PEMPKOWIAK<sup>1</sup>  
JACEK BELDOWSKI<sup>1</sup>  
KSENIA PAZDRO<sup>1</sup>  
ANDRZEJ STANISZEWSKI<sup>1</sup>  
THOMAS LEIPE<sup>2</sup>  
KAY-CHRISTIAN EMEIS<sup>2</sup>

<sup>1</sup>Institute of Oceanology,  
Polish Academy of Sciences,  
Powstańców Warszawy 55, PL-81-712 Sopot, Poland;  
e-mail: pempa@iopan.gda.pl

<sup>2</sup>Baltic Sea Research Institute,  
Seestrasse 15, 18119 Rostock-Warnemünde, Germany

Manuscript received 10 October 2002, reviewed 13 November 2002, accepted 18 November 2002.

**Abstract**

Fluffy Layer Suspended Matter (FLSM) is a layer of fairly concentrated suspended matter resting on the sea floor. Its passage to the depositional basins in the Pomeranian Bay – Arkona Deep System of the Baltic Sea is estimated to take around six months. In the course of this migration, the properties of FLSM change as a result of ageing and the influx of fresh particles from the water column, and possibly also because of mass exchange with the uppermost sediment layers. Measurements of radioisotopes (<sup>210</sup>Po, <sup>210</sup>Pb, <sup>137</sup>Cs) have demonstrated that in shallow water this topmost layer of sediments, from 8 cm to 3 cm in thickness, is subject to mixing. This creates redox profiles favourable to biota and bioturbation.

---

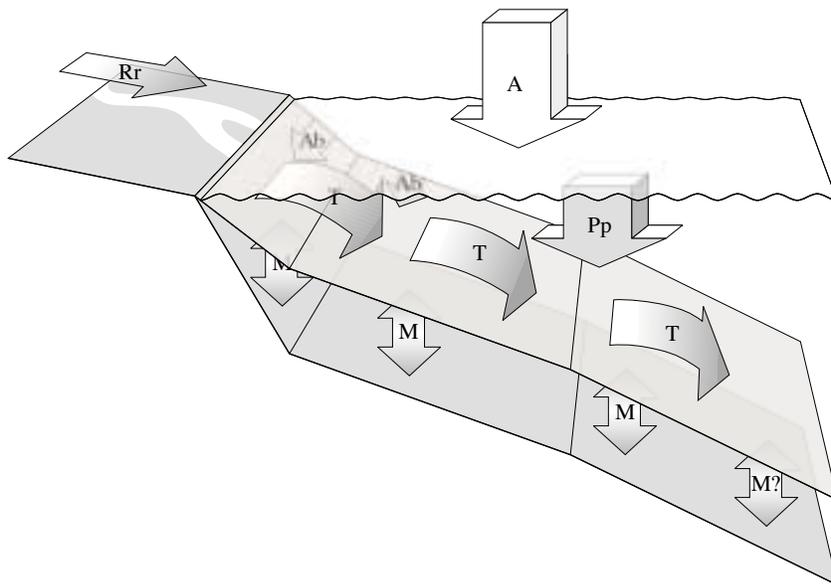
\* The study was supported by the Institute of Oceanology PAS, grant No. 3.1/2002 and the EU FP 5 grant BASYS 3a ‘Coastal to basinal fluxes’.

Basing on  $^{210}\text{Pb}/^{210}\text{Po}$  disequilibria and the  $^{210}\text{Po}$  excess, it was estimated that under steady state conditions from 1.5 to 2.2% of fine fraction (FSF) in the mixed layer of sediments is freshly imported from FLSM. This implies replacement of FSF from the sediments and its incorporation into FLSM. On the assumption that the surface density of FLSM is  $10 \text{ mg cm}^{-2}$ , FSF freshly exported from sediments actually comprises up to 15% of FLSM. Therefore, the properties of FLSM are strongly influenced by the processes taking place in the sediments, although FLSM by definition is independent of sediments.

## 1. Introduction

Material transport across the boundary between the ecosystems of land and sea is important as regards a number of problems ranging from the exposure of biota to pollutants, to elemental biogeochemical cycles. An important aspect in this respect is the formation and transport of suspended matter in the coastal environment. The dispersal of chemical species in the water column is determined by a variety of processes such as coastal circulation, particle absorption and adsorption, vertical and horizontal scavenging. The contribution of the processes to particulate matter formation varies seasonally, since the climate modifies both hydrological factors and biological activity. Numerous studies have been carried out to establish the importance of these factors (Alldredge & Gotschalk 1990, Mayer-Reil 1994, Bale & Morris 1998, Gustafsson et al. 1998, Jago & Jones 1998, Jones et al. 1998, Ransom et al. 1998, Laima et al. 1999, Lund-Hansen et al. 1999, Mayer 1999, Leipe et al. 2000, Löffler et al. 2000, Emeis et al. 2002, Martino et al. 2002).

An important feature of the phenomenon is suspended matter formation and transport in the near-bottom water layer. As presented in Fig. 1, particulate matter there is formed from suspended matter entering the coastal environment with river run-off, produced by abrasion of the shore and shallow sea bottom, primary production, atmospheric deposition, autogenic mineral formation, and anthropogenic discharges. En route from the high-energy, shallow, coastal environment to the low-energy, deep, depositional basins the matter is subject to later resuspension, migration driven by near-bottom currents, and sedimentation (Emeis et al. 2002). Wind-induced currents (Lund-Hansen et al. 1997) and aggregation of suspended matter particles (Thomsen & Ritzrau 1996) are thought to be major factors influencing near-bottom transport. However, in periods of calm weather suspended-matter sedimentation leads to the formation of a high-density fluffy suspended matter (FLSM) layer resting on the sea bottom. Once the FLSM is at rest, both physical and biological processes contribute to the incorporation of the suspended matter into the sediments. The same factors cause fine particles to be released from the sediment



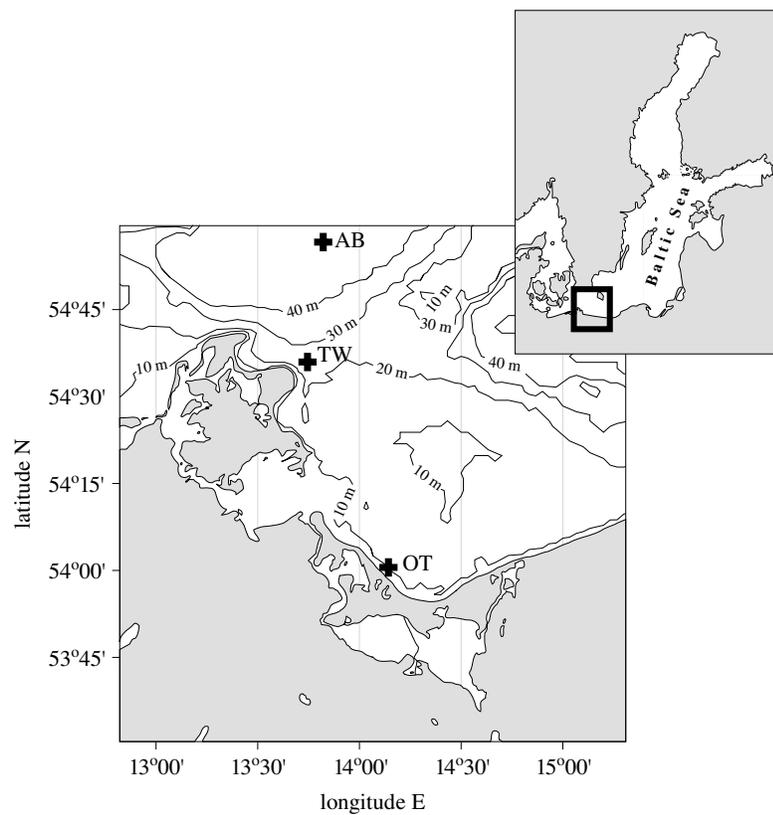
**Fig. 1.** Sketch diagram showing sources and migration of the fluffy layer suspended matter (FLSM) in the coastal environment (A – atmospheric deposition; Rr – river run-off; Ab – abrasion; Pp – primary production; T – resuspension, transport, aggregation, and sedimentation; M – mass exchange between FLSM and the uppermost sediment layer)

and, at least temporarily, incorporated into the suspended matter layer. Being a part of FLSM, these particles share its fate, that is, resuspension, transport, aggregation and sedimentation in the new environment.

This paper presents the results of a study of the processes of mass exchange between the sediment and the near-bottom suspended matter. It was carried out in the Odra (Oder) River mouth – Arkona Deep system in the Baltic Sea. Surface sediments, FLSM, and nepheloid layer suspended matter were collected for the measurement and comparison of properties. Isotopic disequilibrium between  $^{210}\text{Pb}$  and its granddaughter  $^{210}\text{Po}$  was used to indicate the extent of the rapid mass exchange process, while  $^{137}\text{Cs}$  activity profiles were used to validate assumptions concerning past mixing events of the surface sediments. The study was performed in order to quantify the exchange of mass between the uppermost layers of sediments and the overlying suspended matter.

## 2. The study area

The Pomeranian Bay – Arkona Deep system extends from the Odra (Oder) River estuary to the Arkona Basin (Fig. 2). Three sampling stations



**Fig. 2.** The study area (sampling stations: OT – Odas Tonne, TW – Tromper Wiek, AB – Arkona Basin)

**Table 1.** Geographical positions, water depths, and average near-bottom water quality at the sampling stations

Station	Latitude (N)	Longitude (E)	Water depth [m]	Oxygen [ml l <sup>-1</sup> ]	Salinity [PSU]
OT	54°05'	14°10'	16	5.8	7.5
TW	54°36'	13°46'	26	5.9	9.1
AB	54°56'	13°50'	47	2.6	16.4

(Fig. 2) – Odas Tonne (OT), Tromper Wiek (TW), and Arkona Basin (AB) were established there, with the water depth increasing from 16 m at OT to 47 m at AB. The geographical co-ordinates and characteristics of the sampling stations are given in Table 1.

The hydrography of this area is influenced by the Odra with its annual water run-off of about  $17 \text{ km}^3$ , and by the weather conditions, which influence the water circulation. Sediments in the Pomeranian Bay are temporarily covered by a thin fluffy layer of suspended matter (FLSM). On a weight basis the quantity of FLSM is estimated to range from  $50 \text{ g m}^{-2}$  to  $150 \text{ g m}^{-2}$  (Leipe et al. 2000). The layer is most often from 1 cm to 3 cm thick. The sediments are composed of a well sorted fine sand (0.1 to 0.25 mm) grading to muddy sand below 20 m water depth and to mud in the depositional basin of the Arkona Deep.

The wind is the main factor governing the currents in the Pomeranian Bay (Löffler et al. 2000). Density-driven currents and tides are of minor importance. Easterly winds dominate between February and May. At this time surface water moves to the north while the near-bottom water layer is transported to the east. Westerly winds are dominant between June and September, while westerly to south-westerly winds generally occur between October and January. During these periods the surface water is transported eastwards along the coast. This surface current is compensated by a near-bottom current, modified by the bottom topography, that flows towards the Arkona Deep (Siegel et al. 1999). The sampling stations are situated along a transect following the main transport route of the near-bottom water layer and the suspended matter it carries (Leipe et al. 2000). Although the grain size at the sediment surface was different at the sampling stations (Lund-Hansen et al. 1999), the critical shear stress at all stations was surprisingly uniform, corresponding to flow velocities of about  $4 \text{ cm s}^{-1}$ . This low value is due to the presence of an organic-rich layer resting on top of the actual sediment surface – FLSM. A flow velocity of  $4.5 \text{ cm s}^{-1}$  is needed to transport the fluffy material (Leipe et al. 2000). The Arkona Deep acts as a final deposition area for the material, with an average mass accumulation rate from 1.0 to  $1.8 \text{ kg(dry mass) m}^{-2} \text{ year}^{-1}$  (Widrowski & Pempkowiak 1986, Emeis et al. 2000).

### 3. Sources and composition of FLSM

Altogether, the Odra discharges about  $425\,000 \text{ t year}^{-1}$  of suspended solids. The material is rich in organic matter (15.3% particulate organic carbon) and heavy metals (Pb –  $200 \mu\text{g g}^{-1}$ , Zn –  $1700 \mu\text{g g}^{-1}$ , Cu –  $120 \mu\text{g g}^{-1}$ , Cd –  $9 \mu\text{g g}^{-1}$ , Hg –  $2.5 \mu\text{g g}^{-1}$ ) (Emeis et al. 2002). In comparison to pelagic suspended matter (PSM), FLSM is rich in hydrated three-layer and mixed-layer silicates such as illite and smectite (Leipe et al. 2000). This is due to their ability to enhance aggregation and sedimentation of suspended matter particles (Leipe et al. 2000). The organic matter content in FLSM is much smaller than in the nepheloid layer suspended matter

(NLSM), indicating that organic matter may be undergoing mineralisation during sedimentation. This conclusion is supported by the fact that the organic matter content at the deepest station (AB – 47 m) is less than at the shallowest one (OT – 16 m). Fe-oxihydroxides are present in FLSM: the content of these oxides decreases from about 10% at OT to 5% at TW and 3% at AB. Mn-oxihydroxides are present only at OT (2%) (Leipe et al. 2000). This is most likely due to the anoxic conditions in the uppermost sediments, causing reduction of Fe(III) to Fe(II) and of Mn(IV) to Mn(II). The reduced species migrate to the overlying water, where they are oxidised and precipitated (Pempkowiak et al. 2000).

#### 4. Experimental

A rosette sampler connected to a CTD-profiler was used for collecting water samples 5 m above the sea bottom (samples designated as PSM). A bottom-water sampler was used to collect samples (NLSM) 0.4 m above the bottom (Thomsen et al. 1994). The fluffy layer covering the sediment surface was collected by scuba divers, or by a remotely operated vehicle (ROV). The material (FLSM) was siphoned from the sediment surface to an on-board pump and transferred into 100 l containers.

The particulate matter was separated from the water samples by filtration through pre-weighed glass-fibre filters and analysed according to the following procedures. Particulate organic carbon (POC) and particulate organic nitrogen (PON) were measured in a Fission Instruments CHN Analyser. The activity of  $^{210}\text{Po}$  was determined according to the method described in Pempkowiak (1991). In brief, the method consists of the following steps: 0.250 g d.m. of each sample is wet-digested ( $\text{HNO}_3:\text{HClO}_4:\text{HF} = 1:1:3$ ), excess acids are evaporated, the dry residue is redissolved in 0.1 mol  $\text{dm}^{-3}$   $\text{HNO}_3$ , and  $^{210}\text{Po}$  is spontaneously deposited on silver discs. The activity of the deposited  $^{210}\text{Po}$  is measured in a 1024 Multichannel Analyser (Polon) coupled with a Si surface barrier detector (Canberra). In the present study, the yield measured by  $^{209}\text{Po}$  (Amersham) addition prior to digestion ranged from 84.3% to 106.1% ( $92.4 \pm 4.2\%$ ). Repeated measurements of  $^{210}\text{Po}$  in the samples after 12 months of in-growth following the initial extraction of  $^{210}\text{Po}$  were used for the  $^{210}\text{Pb}$  activity measurements. Excessive activity of  $^{210}\text{Po}$  was obtained subtracting ‘repeated’  $^{210}\text{Po}$  activity from the ‘initial’  $^{210}\text{Po}$  activity.

Sediment accumulation rates at station AB were determined using the constant rate of supply model (Robbins 1978, Pempkowiak 1991).

$^{137}\text{Cs}$  was determined by  $\gamma$ -spectrometry with a HPGe detector, with an energy resolution of 1.8 keV for  $^{60}\text{Co}$  (1332 keV), and a relative efficiency of 30%.

## 5. Results and discussion

The results for samples collected at station OT are presented in Table 2, for TW in Table 3 and for AB in Table 4.

**Table 2.** Characteristics of the samples collected at station OT

SAMPLE*	C <sub>org</sub> [%]	$\delta^{13}\text{C}$ [‰]	$\delta^{15}\text{N}$ [‰]	$^{137}\text{Cs}$ [mBq g <sup>-1</sup> ]	$^{210}\text{Pb}$ [mBq g <sup>-1</sup> ]	$^{210}\text{Po}_{\text{ex}}$ [mBq g <sup>-1</sup> ]	Concentration [mg dm <sup>-3</sup> ]
PSM	22.6		9.6	nd**	nd	nd	2.1
NLSM	14.1		9.1	nd	62.8	32.8	3.5
FLSM	8.7	-25.0	9.9	89.7	61.5	26.5	10 × 10 <sup>3</sup>
S 0–2 cm	7.2		9.8	56.2	38.7	1.4	
S 2–4 cm	7.3		nd	57.4	39.4	2.3	
S 4–6 cm	6.9		nd	53.5	36.1	bdl***	
S 6–8 cm	7.0		nd	54.6	38.3	bdl	
S 8–10 cm	6.1		nd	2.7	30.4	bdl	
S 10–12 cm	6.4		nd	bdl	31.5	bdl	

\* PSM – pelagic (5 m above bottom) suspended matter,

NLSM – nepheloid layer (40 cm above bottom) suspended matter,

FLSM – fluffy layer (0–3 cm above bottom) suspended matter,

S – fine fraction (< 0.67 mm) retrieved from 2 cm thick slices of a sediment core,

\*\* not determined,

\*\*\* below detection limit.

**Table 3.** Characteristics of the samples collected at station TW

SAMPLE*	C <sub>org</sub> [%]	$\delta^{13}\text{C}$ [‰]	$\delta^{15}\text{N}$ [‰]	$^{137}\text{Cs}$ [mBq g <sup>-1</sup> ]	$^{210}\text{Pb}$ [mBq g <sup>-1</sup> ]	$^{210}\text{Po}_{\text{ex}}$ [mBq g <sup>-1</sup> ]	Concentration [mg dm <sup>-3</sup> ]
PSM	28.6		6.1	nd**	nd	nd	1.9
NLSM	7.2		6.4	nd	nd	nd	5.9
FLSM	6.1	-24.0	7.0	94.6	114.7	28.6	10 × 10 <sup>3</sup>
S 0–2 cm	6.6		6.8	85.4	104.2	1.2	
S 2–4 cm	6.2		nd	76.2	104.1	1.3	
S 4–6 cm	5.6		nd	75.8	102.9	0.9	
S 6–8 cm	5.7		nd	1.7	32.3	bdl***	
S 8–10 cm	5.4		nd	bdl	33.5	bdl	
S 10–12 cm	5.5		nd	bdl	31.6	bdl	

Symbols and explanations as in Table 2.

**Table 4.** Characteristics of the samples collected at station AB

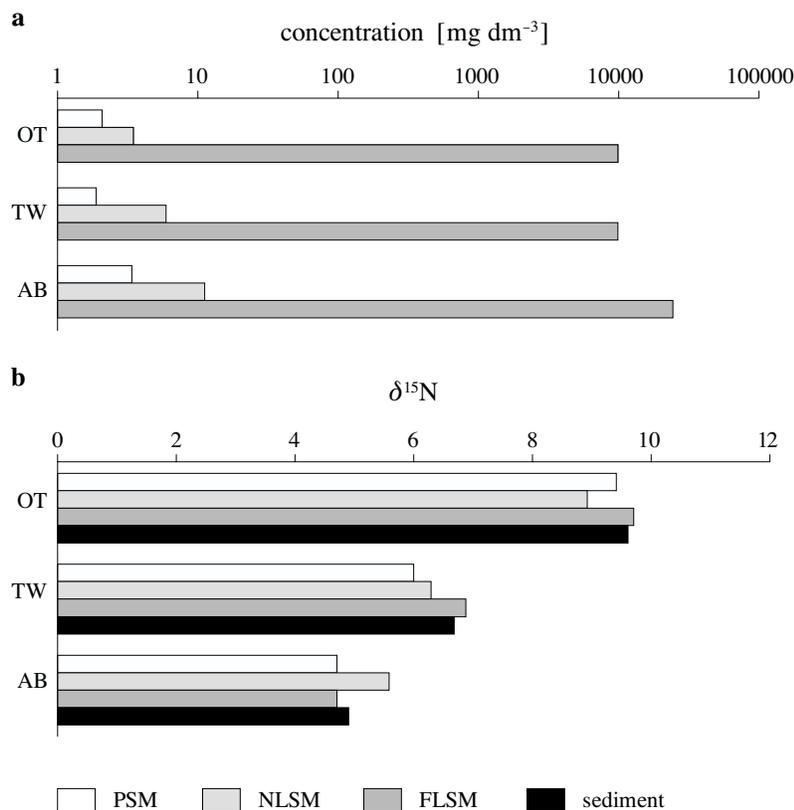
SAMPLE*	C <sub>org</sub>	$\delta^{13}\text{C}$	$\delta^{15}\text{N}$	$^{137}\text{Cs}$	$^{210}\text{Pb}$	$^{210}\text{Po}_{\text{ex}}$	Concentration
	[%]	[‰]	[‰]	[mBq g <sup>-1</sup> ]	[mBq g <sup>-1</sup> ]	[mBq g <sup>-1</sup> ]	[mg dm <sup>-3</sup> ]
PSM	7.6		4.8	nd**	nd	nd	3.4
NLSM	5.1		5.7	nd	nd	nd	11.2
FLSM	6.4	-22.1	4.8	107.5	137.4	6.2	$2.5 \times 10^4$
S 0–2 cm	6.4		5.0	100.1	211.5	bdl***	
S 2–4 cm	6.0		nd	94.3	130.9	bdl	
S 4–6 cm	5.8		nd	64.2	97.0	bdl	
S 6–8 cm	5.4		nd	1.8	69.4	bdl	
S 8–10 cm	5.3		nd	1.3	61.3	bdl	
S 10–12 cm	5.6		nd	bdl	55.2	bdl	
S 12–14 cm	nd		nd	bdl	52.1		
S 16–18 cm	nd		nd	bdl	51.3		
S 18–20 cm	nd		nd	nd	51.9		

Symbols and explanations as in Table 2.

It follows from the data that particulate matter is progressively more abundant in the water column towards the sea floor (Fig. 3a). There is also an increase in particulate matter in the nepheloid layer in the transect from OT to AB. The latter phenomenon is caused by frequent resuspension events in the high-energy shallow environment of OT as compared to the environment of both the deeper stations, favouring a more uniform distribution of particulate matter at the shallow stations. The same applies to the steeper concentration gradient of particulate matter at AB as compared to OT. The slightly larger concentration of particulate matter in the surface water (PSM) at OT is caused by the spring algal bloom there. The high proportion of organic matter in PSM at OT corroborates this conclusion. There is a decrease in  $\delta^{13}\text{C}$  towards station AB, indicating the increasing contribution of autochthonous organic matter to FLSM. This is caused by both mineralisation of allochthonous organic matter and its dilution in the marine environment as a result of primary production. Both vertical and horizontal gradients of  $\delta^{15}\text{N}$  substantiate this conclusion (Fig. 3b).

Vertical profiles of the measured radionuclides are presented in Figs. 4a, 4b, and 4c respectively for stations OT, TW, and AB.

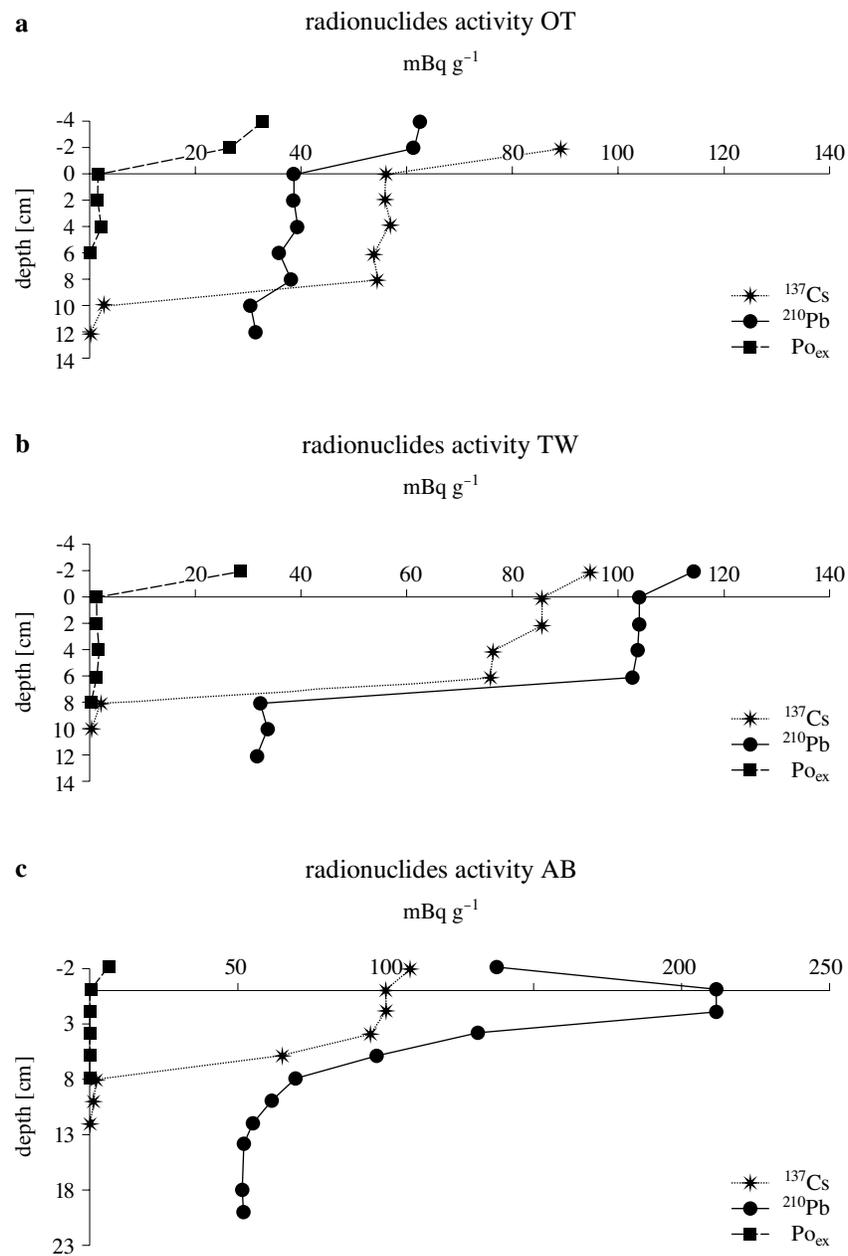
The  $^{210}\text{Pb}$  activity profile at OT exhibits a rapid activity shift 8 cm below the sediment-water interface. Below this depth only supported  $^{210}\text{Pb}$



**Fig. 3.** Concentrations of particulate matter (a) and  $\delta^{15}\text{N}$  (b) in the water column (PSM – 5 m above sea floor; NLSM – 0.4 m above sea floor; FLSM – 0–3 cm above sea floor, and the uppermost (0–2 cm) sediment layer at the sampling stations OT, TW, AB)

occurs in the sediment column. The sharp increase in  $^{210}\text{Pb}$  activity in the 6–8 cm sediment layer is due to the presence of supported  $^{210}\text{Pb}$ . The constant activity profile in sediments above 8 cm indicates that the activity shift was caused by a rapid mixing event. Both bioturbation and mechanical mixing due to wave action may explain this shift since the area abounds in macrofauna (Warzocha 1994), whereas storms causing resuspension of the bottom sediments occur on average once in 12 years (Emeis et al. 2002). The  $^{137}\text{Cs}$  profile provides confirmation of such an explanation: it shows constant activities down to the 6–8 cm layer. A slight increase in  $^{210}\text{Pb}$  activity in the 0–4 cm layer of sediments may have been caused by recent rapid mixing.

The mixing process at the FLSM/sediment interface could be followed using  $^{234}\text{Th}$  (Tanaka & Tsunogai 1983, Tanaka et al. 1983, Gustafsson et al.



**Fig. 4.** Activity profiles of <sup>137</sup>Cs, <sup>210</sup>Pb and <sup>210</sup>Po<sub>ex</sub> in the NLSM, FLSM, and sediments at (a) station OT, (b) station TW, and (c) station AB

1998). In this study <sup>210</sup>Po/<sup>210</sup>Pb isotopic disequilibrium was used because of the simplicity and precision of the measurements. In the southern Baltic this disequilibrium was first reported by Skwarzec (1995).

In this study, the presence of  $^{210}\text{Po}_{\text{excess}}$  was found in NLSM. It is an example of the polonium and lead isotopic disequilibrium frequently found in the marine environment (Poet et al. 1972, Bacon et al. 1976, Nozaki & Tsunogai 1976, Schell 1977, Heyraud & Cherry 1983, Skwarzec 1995). The disequilibrium is apparently caused by preferential accumulation of polonium by biota (Tanaka & Tsunogai 1983, Tanaka et al. 1983, McDonald et al. 1986, Skwarzec & Bojanowski 1988, Skwarzec & Falkowski 1988, Skwarzec 1995). The excess of  $^{210}\text{Po}$  in the uppermost sediment layer (Tables 2–4) is caused by the rapid entry of fine particles from the FLSM to the uppermost sediments. The mechanism leading to the  $^{210}\text{Po}$  and  $^{210}\text{Pb}$  activity profiles may be as follows: scavenging of  $^{210}\text{Po}$  from the water column (Schell 1977, Heyraud & Cherry 1983) and accumulation by macrobenthos (McDonald et al. 1986, Skwarzec 1995) leads to isotopic disequilibrium with  $^{210}\text{Pb}$  in FLSM (Skwarzec 1995). FLSM is utilised by benthic organisms, thus part of it is incorporated into the sediments. The process must be a rapid one, since the  $^{210}\text{Pb}/^{210}\text{Po}$  ratios in FLSM are similar to those in the uppermost sediment layer.

An isotopic disequilibrium between  $^{210}\text{Po}$  and  $^{210}\text{Pb}$  in FLSM was also recorded in the FLSM at station TW. It was caused by the same phenomenon as at OT. The  $^{210}\text{Pb}$  activity in the uppermost, 6 cm thick, layer of the sediment column is greater than that at OT because of the higher proportion of the fine fraction in the sediments there (Leipe et al. 2000). The unsupported  $^{210}\text{Pb}$  reaches as deep into the sediments as the excess  $^{210}\text{Po}$ . This indicates that the uppermost layer has been subjected to rapid recent mixing, most likely as a result of bioturbation. There is no difference between the depths where  $^{137}\text{Cs}$ , the unsupported  $^{210}\text{Pb}$  profile, and the excess  $^{210}\text{Po}$  profile occur, an indication that no processes other than bioturbation leading to surface mixing have taken place. This inference is further corroborated by the slight increase of  $C_{\text{org}}$  in the uppermost 6 cm thick layer of sediments. Below 6 cm down to the sediments the activity of  $^{210}\text{Pb}$  is steady. It is regarded as supported  $^{210}\text{Pb}$  profile, since no  $^{137}\text{Cs}$  is present there.

The sediment core at site AB represents the muddy sediments typical of the Arkona Deep. Reducing conditions ( $E_{\text{h}} = -150$  mV) exist in the Arkona Deep sediments. The exponential decrease in  $^{210}\text{Pb}$  activity to  $51$  mBq  $\text{g}^{-1}$  at a depth of 20 cm below the sediment-water interface indicates steady sedimentation and no surface mixing. The lack of isotopic  $^{210}\text{Pb}/^{210}\text{Po}$  disequilibrium indicates that no rapid recent mixing has occurred. Bioturbation has not taken place because of the reducing conditions in the sediment column, while the considerable depth prevents mixing due to wave action. The sedimentation rate calculated from the  $^{210}\text{Pb}$  profile is

$0.91 \pm 0.02 \text{ mm year}^{-1}$ . The rate is similar to the earlier findings in the area (Pempkowiak 1991). The depth of actual  $^{137}\text{Cs}$  activity (4–6 cm) is the same as the one calculated using the sedimentation rate derived from the radiolead profile (5 cm). This validates the calculated sedimentation rate. The  $^{137}\text{Cs}$  maximum, often found in undisturbed sediments (McCall & Sørensen 1984, Wong & May 1984) is absent owing to the delayed river run-off maximum (Müller et al. 1980, Widrowski & Pempkowiak 1986).

## 6. Mass balance

The isotopic disequilibrium found in FLSM and in the uppermost layer of sediments combined with the concentration of particulate matter in FLSM and the proportion of fine sediment fraction (FSF) are used to calculate mass exchange between sediments and water. It is based on the assumption that the exchange is instantaneous once the FLSM is formed on the sandy bottom of the study area. The contribution of FLSM to FSF is calculated from the following equation:

$$C = \frac{{}^{210}\text{Po}_{\text{excess}}(\text{FLSM}) \times L_{(\text{FLSM})}}{\sum {}^{210}\text{Po}_{\text{excess}}(\text{FSF}) \times D \times F \times d} 100 \quad [\%], \quad (1)$$

where  ${}^{210}\text{Po}_{\text{excess}}$  is the excess activity of  ${}^{210}\text{Po}$  in particulate matter (FLSM), or sediments (FSF);  $L_{(\text{FLSM})}$  is the surface load of FLSM;  $D$  is depth of  ${}^{210}\text{Po}_{\text{excess}}$ ;  $F$  is the proportion of the fine sediment fraction;  $C$  is the contribution of FLSM to FSF;  $d$  is the density of sediments.

The contribution of FSF to FLSM ( $S$ ) is

$$S = C \times \frac{L_{(\text{FSF})}}{L_{(\text{FLSM})}} = \frac{D \times F \times d}{L_{(\text{FSF})}} \quad [\%], \quad (2)$$

where  $L_{(\text{FSF})}$  is the cumulative surface load of FSF.

Taking the surface FLSM load and values of  $D$  from Tables 3 and 4, and assuming  $F = 0.006$  (OT) or  $F = 0.025$  (TW) and  $d = 2.5 \text{ g cm}^{-3}$ , the respective contributions of FLSM to FSF under steady state conditions are 4.5% and 1.7% at stations OT and TW, while the FSF contributions to FLSM are 12.5% at OT and 3.6% at TW.

## References

- Allredge A., Gotschalk S., 1990, *The relative contribution of marine snow of different origins to biological processes in coastal waters*, Cont. Shelf Res., 10, 41–58.
- Bacon M., Spencer D., Brewer P., 1976,  *${}^{210}\text{Pb}/{}^{226}\text{Ra}$  and  ${}^{210}\text{Po}/{}^{210}\text{Pb}$  disequilibria in seawater and suspended particulate matter*, Earth Planet. Sci. Lett., 32, 277–296.

- Bale A., Morris A., 1998, *Organic carbon in suspended particulate material in the North Sea: Effect of mixing resuspended and background particles*, Cont. Shelf Res., 18, 1333–1345.
- Emeis K.-C., Christiansen C., Edelvang K., Jähmlich S., Kozuch J., Laima M., Leipe T., Löffler A., Lund-Hansen L., Miltner A., Pazdro K., Pempkowiak J., Shimmiel G., Shimmiel T., Smith J., Voss H., Witt G., 2002, *Material transport from the nearshore to the basinal environment in the southern Baltic Sea. II: Origin and properties of material*, J. Mar. Sys., 35, 151–168.
- Emeis K.-C., Struck U., Leipe T., Pollehne F., Kunzendorf H., Christiansen C., 2000, *Changes in the C, N, P burial rates in some Baltic Sea sediments over the last 150 years – relevance to P regeneration rates and the phosphorus cycle*, Mar. Geol., 167, 43–59.
- Gustafsson O., Buesseler K., Geyer W., Moran S., Gschwend P., 1998, *An assessment of the relative importance of horizontal and vertical transport of particle-reactive chemicals in the coastal ocean*, Cont. Shelf Res., 18, 805–829.
- Heyraud M., Cherry R., 1983, *Correlation of  $^{210}\text{Po}$  and  $^{210}\text{Pb}$  enrichments in the sea-surface microlayer with neuston biomass*, Cont. Sea Res., 1, 283–293.
- Jago C., Jones S., 1998, *Observation and modelling of the dynamics of benthic fluff resuspended from a sandy bed in the southern North Sea*, Cont. Shelf Res., 18, 1255–1282.
- Jones S., Jago C., Bale A., Chapman D., Howland R., Jackson J., 1998, *Aggregation and resuspension of suspended particulate matter at a seasonally stratified site in the southern North Sea: Physical and biological controls*, Cont. Shelf Res., 18, 1283–1309.
- Laima M., Lund-Hansen L., Pazdro K., Christiansen C., Emeis K.-C., 1999, *Near bottom fluxes and composition of suspended matter in the Pomeranian Bay*, Oceanologia, 41 (3), 335–353.
- Leipe T., Loeffler A., Emeis K.-C., Jaehmlich S., Bahlo R., Siervogel K., 2000, *Vertical patterns of suspended matter characteristics along a coastal-basin transect in the Western Baltic Sea*, Estuar. Coast. Shelf Sci., 51, 789–804.
- Löffler A., Leipe T., Emeis K.-C., 2000, *The fluffy layer in the Pomeranian Bight (western Baltic Sea): geochemistry, mineralogy and environmental aspect*, Meyniana, 52, 85–100.
- Lund-Hansen L., Christiansen C., Jensen O., Laima M., 1999, *The LABREX chamber for studying the critical shear stress for fine-grained sediments*, Danish J. Geogr., 99, 1–7.
- Lund-Hansen L., Valer J., Pejrup H., Jensen A., 1997, *Sediment fluxes, resuspension and accumulation rates in two wind-exposed coastal sites and in a sheltered bay*, Estuar. Shelf Sci., 44, 521–531.
- Martino M., Turner A., Niemmo M., Millward G., 2002, *Resuspension, reactivity and recycling of trace metals in the Mersey Estuary*, Mar. Chem., 77, 171–186.
- Mayer L., 1999, *Extent of coverage of mineral sources by organic matter in marine sediments*, Geochim. Cosmochim. Acta, 63, 207–215.

- Mayer-Reil L., 1994, *Microbial life in sedimentary biofilms – the challenge to microbial ecologists*, Mar. Ecol. Prog. Ser., 112, 303–311.
- McCall P., Sørensen J., 1984,  $^{137}\text{Cs}$  and  $^{210}\text{Pb}$  transport and geochronologies in urbanized reservoirs with rapidly increasing sedimentation rates, Chem. Geol., 44, 33–65.
- McDonald P., Fowler S., Heyrand M., Baxter M., 1986, *Polonium-210 in mussels and its implications for environmental alpha-autoradiography*, J. Environ. Radioact., 3, 293–303.
- Müller G., Dominik J., Reuther R., Malisch R., Schultz E., Acker L., 1980, *Sedimentary record of environmental pollution in the Western Baltic Sea*, Naturwissenschaften, 67, 595–600.
- Nozaki Y., Tsunogai S., 1976,  $^{226}\text{Ra}$ ,  $^{210}\text{Pb}$  and  $^{210}\text{Po}$  disequilibria in the western North Pacific, Earth Planet. Sci. Lett., 32, 313–321.
- Pempkowiak J., 1991, *Enrichment factors of heavy metals in the southern Baltic surface sediments dated with  $^{210}\text{Pb}$  and  $^{137}\text{Cs}$* , Environ. Int., 17, 421–428.
- Pempkowiak J., Chiffoleau J.-F., Stanisławski A., 2000, *The vertical and horizontal distribution of selected trace metals in the Baltic Sea off Poland*, Estuar. Coast. Shelf Sci., 51, 115–125.
- Pempkowiak J., Kożuch J., Szymelfenig M., 1995, *Meiobenthic organisms as indicators of mixing processes in the Baltic surface sediments*, PSNI Mar. Ecol., 17 (1)–(3), 175–178.
- Poet S., Moore H., Martel E., 1972, *Lead-210, bismuth-210, and polonium-210 in the atmosphere: accurate ratio measurement and application to aerosol residue time determination*, J. Geophys. Res., 77, 6515–6527.
- Ransom B., Kim D., Kastner M., Wainwright S., 1998, *Organic matter preservation on continental slopes: Importance of mineralogy and surface area*, Geochim. Cosmochim. Acta, 62, 1329–1345.
- Robbins J., 1978, *Geochemical and geophysical application of radioactive lead*, [in:] *The biogeochemistry of lead in the environment*, J. O. Nriagu (ed.), Elsevier, Amsterdam, 285–393.
- Schell W., 1977, *Concentrations, physico-chemical states and near residence times of  $^{210}\text{Pb}$  and  $^{210}\text{Po}$  in marine and estuarine waters*, Geochim. Cosmochim. Acta, 41, 1019–1031.
- Siegel H., Gerth M., Mutzke A., 1999, *Dynamics of the Oder river plume in the Southern Baltic Sea: satellite data and numerical modelling*, Cont. Shelf Res., 19, 1143–1159.
- Skwarzec B., 1995, *Polonium, uranium and plutonium in the ecosystem of the southern Baltic Sea*, Rozpr. i monogr., Inst. Oceanol. PAN, Sopot, 6, 179 pp., (in Polish).
- Skwarzec B., Bojanowski R., 1988, *The  $^{210}\text{Po}$  content in sea water and its accumulation in southern Baltic plankton*, Mar. Biol., 97, 301–307.
- Skwarzec B., Falkowski L., 1988, *Accumulation of  $^{210}\text{Po}$  in Baltic invertebrates*, J. Environ. Radioact., 8, 99–109.

- Tanaka N., Takeda Y., Tsunogai S., 1983, *Biological effect of removal of Th-234, Po-210 and Pb-210 from surface water in Funaka Bay, Japan*, Geochim. Cosmochim. Acta, 47, 1783–1790.
- Tanaka N., Tsunogai S., 1983, *Behaviour of  $^7\text{Be}$  in Funaka Bay, Japan with reference to those of insoluble nuclides,  $^{234}\text{Th}$ ,  $^{210}\text{Po}$  and  $^{210}\text{Pb}$* , Geochim. J., 17, 9–17.
- Thomsen L., Graf G., Hartens V., Steen E., 1994, *An instrument for sampling water from the benthic boundary layer*, Cont. Shelf Res., 14, 77–91.
- Thomsen L., Ritzrau W., 1996, *Aggregate studies in the benthic boundary layer at a continental margin*, J. Sea Res., 36, 143–146.
- Warzocha J., 1994, *Spatial distribution of macrofauna in the Baltic*, Bull. Sea Fish. Inst., 1 (131), 47–60.
- Widrowski H., Pempkowiak J., 1986, *The history of surface sediments in the Southern Baltic*, Proc. 15th Conf. Baltic Oceanogr., Mar. Pollut. Lab., Copenhagen, 656–671.
- Wong G., May C., 1984, *Caesium-137, metals and organic carbon in the sediments of the James River estuary, VA*, Estuar. Coast. Shelf Sci., 18, 37–49.